

Energy Budget of Nonlinear Internal Waves near Dongsha

Ren-Chieh Lien

Applied Physics Laboratory
University of Washington
Seattle, Washington 98105

phone: (206) 685-1079 fax: (206) 543-6785 email: lien@apl.washington.edu

Eric A. D'Asaro

Applied Physics Laboratory
University of Washington
Seattle, Washington 98105

phone: (206) 685-2982 fax: (206) 543-6785 email: dasaro@apl.washington.edu

Award Number: N00014-05-1-0284

LONG-TERM GOALS

Our long-term scientific goal is to understand the mechanisms by which mixing occurs in the ocean and thereby help develop improved parameterizations of mixing for ocean models. Mixing within the stratified ocean is our particular focus as the complex interplay of internal waves from a variety of sources and turbulence makes this a current locus of uncertainty. In this study, our broad focus is on the energy sources of nonlinear internal waves (NLIWs) in a complex environment of strong internal tides and abrupt topography (continental shelf and slope). We expect a rapid evolution of internal tides and NLIWs, and we aim to understand their dynamics, energy cascade, and role in mixing.

OBJECTIVES

The primary objectives of the present project are 1) to identify the generation site and understand the generation mechanism of NLIWs, 2) to understand the evolution of NLIWs interacting with abrupt topography, 3) to quantify the energy budget and energy cascade from internal tides to NLIWs, and 4) to quantify the seasonal variation of the energy of NLIWs near Dongsha Island in the northern South China Sea (SCS). Our particular interest is to understand the energy cascade from barotropic tides, internal tides, NLIWs, to turbulence mixing in the northern SCS, and to understand the evolution of NLIWs interacting with the shoaling continental slope.

APPROACH

Our approach is to take direct observations of NLIWs near Dongsha Island where most of NLIW activities were captured in satellite images. Primary platforms include an ADV Lagrangian Float, a bottom mounted ADCP mooring array, shipboard EK500, marine radar, ADCP, and CTD. Our main goals are to quantify the energy budget and evolution of NLIWs across the rapid shoaling continental slope and the gentle plateau near Dongsha Island and to exploit and quantify the seasonal variation of the characteristics of NLIWs.

WORK COMPLETED

We have completed three components of this observational experiment: (1) pilot observations in 2005, (2) extended observations in 2006-2007, and (3) intensive observations in 2007.

Pilot Observations (April 2005)

In April 2005, we conducted a two-week observational experiment near Dongsha. Large-amplitude NLIWs, greater than 150 m, and strong turbulence mixing were observed by the Lagrangian float and shipboard sensors including ADCP, CTD, EK500, and X-band marine radar. Combined remote sensing and in-situ measurements provided detailed properties of large-amplitude NLIWs.

Extended ADCP Observations (June 2006-May 2007)

An array of three ADCP moorings was deployed along the prevailing path of NLIWs near Dongsha in June 2006. Three ADCPs were serviced once and recovered in May 2005. Two of three moorings took velocity measurements for about 11 months. These long-term observations of NLIWs allow us to (1) quantify the seasonal variation of NLIW energy, (2) map the geographical distribution, (3) better understand the dynamics of NLIW evolution over the shoaling topography, and (4) assess the model prediction skill of NLIWs.

Intensive Observations (April-May 2007)

We participated in the multi-ships intensive observation experiment near Dongsha in April-May 2007. Our main goals were to understand the interaction of internal waves, including internal tides and NLIWs, with the rapid shoaling continental slope, and to quantify the energy budget of internal waves near Dongsha. Our primary instruments included Scripps' fast CTD profiler, and shipboard ADCP, CTD, EK500, and marine radar. During the cruise, two McLane moored profilers were deployed, one on the continental slope and the other on the Dongsha plateau (Fig. 2).

RESULTS

Bio-Physical Coupling: Pilot Whales Following NLIWs in SCS

Schools of pilot whales were observed behind the large-amplitude NLIWs. Moore and Lien (2007) proposed that prey availability for pilot whales in the SCS is influenced by the aggregative properties of NLIWs. This is the first report of cetaceans associated with internal waves. We expect that the ecosystem in the SCS is strongly modulated by properties of NLIWs.

Shoaling NLIW: Trapped Core and Fission Process

In April 2005, large-amplitude NLIWs with the vertical displacement of more than 150 m appeared diurnally during the spring tide (Fig. 3). Within NLIWs, we found ~75-m of turbulent overturning (Fig. 4b). Trains of NLIWs appeared on the Dongsha plateau and evolved rapidly as they propagated westward into the shallower water (Fig. 4a).

On April 28, we followed an NLIW with the vertical displacement of ~170 m propagating westward from the upper continental slope (Fig. 5). At the third encounter, a second wave appeared behind the

leading NLIW. The second NLIW had a smaller vertical displacement and propagated slower than the leading NLIW.

The NLIW evolved rapidly while propagating up the continental slope (Fig. 6). At the first encounter, the maximum westward current velocity U_{max} was nearly equal to the propagation speed C , 2 m s^{-1} . Propagating up onto the continental slope, the NLIW reduced its wave speed dramatically from 2 ms^{-1} to 1.2 m s^{-1} in 3.5 hrs within 20 km, whereas the maximum westward current velocity remained at 2 ms^{-1} . The Froude number, $Fr = U_{max} / C$, was greater than one, suggesting the presence of a re-circulating core within the NLIW, confirmed by the shipboard ADCP velocity and the derived stream lines (Fig. 7). The trapped core had a $\sim 100\text{-m}$ vertical scale, and $\sim 500\text{-m}$ horizontal scale containing two counter rotating vortices (closed streamlines).

The zonal velocity averaging within the trapped core should be equal to the wave propagation speed. Assuming the meridional extent of NLIW of 10 km, the net mass transport of the trapped core is $\sim 2 \text{ Sv}$. The trapped core propagates for $\sim 30 \text{ km}$ (Fig. 6) and causes significant transient mass transport in the SCS. The development of the trailing wave is shown in Fig. 7. We hypothesize that this fission process is triggered by the dissipative trapped core. Because of the strong turbulence mixing, water particles within the trapped core fail to make the complete loop, fall behind the trapped core, and form the second wave.

We compare properties of observed NLIWs with solutions of the fully nonlinear steady-state non-dissipative internal solitary wave model described by Dubriel-Jacotin-Long (DJL) (Fig. 8). The NLIW observed at the first encounter, before the formation of the trapped core, agrees well with the DJL solution. Observations taken after the first encounter do not agree with the DJL solutions. This is not unexpected because of the strong unsteadiness and dissipation within the NLIW. Interestingly, all observations during the first five encounters agree with the DJL solution at the first encounter. Collaborations with small-scale numerical modelers will be necessary to understand the dynamics of the shoaling NLIWs, the trapped core formation, and the fission process.

Radar Detection of NLIW Properties

Surface waves are modulated by NLIWs. In SCS, the interaction is extremely strong due to the large horizontal velocity convergence of NLIWs (Chang et al., 2007). Photos taken during the experiment show a sequence of calm sea surface, breaking surface waves, isotropic surface waves, and boils during the passage of the NLIW during a calm day (Fig. 9). The shipboard marine radar reveals the strongest scattering strength corresponding to the strongest horizontal velocity convergence measured by the shipboard ADCP (Fig. 9b and 9c).

A composite view of the sea surface scattering strength, the horizontal velocity convergence and the vertical displacement of NLIWs is constructed (Fig. 10). The strongest surface scattering strength appears at the front portion of the NLIW, $\sim 0.5 \lambda_{\eta/2}$, where $\lambda_{\eta/2}$ is the wave-width at half of the maximum amplitude.

An empirical formula between the surface scattering strength and horizontal velocity convergence is obtained (Fig. 11). It shows a linear relation at low horizontal velocity convergence, and reaches an asymptotic value at high horizontal velocity convergence, indicating a saturation state presumably due to breaking surface waves. Following the power-law and the linear fits between the scattering strength and the local wind speed, the maximum surface scattering enhancement by NLIWs is equivalent to that

caused by a wind of $\sim 6 \text{ m s}^{-1}$ with surface waves of $\sim 1.5 \text{ m}$, according to the Beaufort wind scale. In other words, the surface scattering enhancement induced by NLIWs will be overwhelmed by that induced by the wind and can not be identified with the X-band radar when the wind is stronger than 6 m s^{-1} . The vertical displacement of NLIWs is found to be proportional to the spatial integration of the surface scattering strength (not shown). Our analysis concludes that in the low-wind condition remote sensing measurements may provide useful predictions of horizontal velocity convergences, amplitudes, and spatial structures of NLIWs. Further applications and modification of our empirical formulas in different conditions of wind speed, surface waves, and NLIWs, or with other remote sensing methods should be investigated.

Rapid Evolution of Shoaling NLIWs Revealed by Mooring Array

Our long-term moored ADCP velocity observations reveal complex patterns of the energy evolution across three mooring sites (illustrated in Fig. 12). These examples show (1) a monotonic decrease of NLIW energy from LR1 to LR3, (2) an increase of NLIW energy at LR2, and (3) an increase of NLIW energy at LR3. Cases (1) and (2) are typical. The fourth case shows the development of a wave train between LR2 and LR3.

At LR1, the maximum westward horizontal velocity U_{max} within the NLIW is always smaller than the propagation speed C , suggesting the absence of a trapped core (Fig. 15). At LR2, nearly 40% of NLIWs has $U_{max} > C$, suggesting the presence of trapped cores. Apparently, trapped cores are formed when NLIWs propagate from LR1 to LR2, within a short distance of ~ 9 miles.

Seasonal Variations of NLIW Properties

Long-term moored observations reveal seasonal variations of the arrival time (Fig. 16), the propagation speed (Fig. 17), and the total kinetic energy of NLIWs (Fig. 18). In spring, NLIWs arrive at LR1 nearly the same time every day and once a day, except occasional weaker NLIWs in the early morning, consistent with observations in the pilot experiment. In Jan.-Feb. 2007, only a few weak NLIWs appear. During the rest of the year, two distinct groups of NLIWs arrive at the mooring site at different times of the day. From mid June to November, the arrival time shows a ~ 40 -min delay every day.

Averaging over ~ 1 -year of observations, the total kinetic energy of NLIWs decreases westward from LR1 to LR3 (Fig. 13a). Comparing the total kinetic energy of individual waves from LR1 to LR3 (Fig. 14) suggests that (1) strong NLIWs at LR1 lose energy before arriving at LR2, (2) weak NLIWs at LR1 gain energy before arriving at LR2, and (3) NLIWs mostly lose energy when propagating from LR2 to LR3. Note that the loss of energy may be due to either the dissipation process or the fission process.

Generation Region of Trapped-Core NLIWs

The rapid decrease of the propagation speed of NLIWs on the shoaling continental slope may lead to the formation of the trapped core, as observed in the 2005 pilot experiment (Fig. 6). Our long-term moored ADCP measurements show that the average propagation speed of NLIWs decreases from 2 m s^{-1} at LR1, 1.8 m s^{-1} at LR2, to 1.3 m s^{-1} at LR3 (Fig. 13), consistent with observations in the pilot experiment (Fig. 6).

The propagation speed of NLIWs increases from June to Oct. 2006, drops in Nov. 2006, and remains at a low speed thereafter (Fig. 17). This pattern is robust at all three mooring sites. The vertical kinetic

energy of NLIWs has a general decreasing trend from June to Dec. 2006 (Fig. 18), is the weakest in Jan.-Feb. 2007, and increases in spring of 2007.

IMPACT/APPLICATION

Our analysis concludes that NLIWs evolve rapidly across the upper flank of the continental slope and the Dongsha plateau via complicated processes, e.g., the formation of trapped cores, and the development of wave trains. These processes are responsible for the strong dissipation of NLIWs in SCS. Our analysis of combined remote sensing and in-situ measurements yields a model for predicting NLIW properties applicable to satellite observations. Further long-term observations of NLIWs in the vicinity of Dongsha plateau are in progress and will provide a better prediction of NLIWs in SCS.

REFERENCES

Chang, M.-H., R.-C. Lien, T. Y. Tang, Y. J. Yang, and J. Wang, Surface signatures of nonlinear internal waves in South China Sea: properties and predictions, *J. Atmos. Ocean. Tech.*, accepted, 2007.

Chang, M.-H., R.-C. Lien, T. Y. Tang, E. A. D'Asaro, and Y. J. Yang, Energy flux of nonlinear internal waves in northern South China Sea, *Geophys. Res. Lett.*, **33**, L03607, doi:10.1029/2005GL025196, 2006.

Lien, R.-C., T. Y. Tang, M. H. Chang, E. A. D'Asaro, Energy of nonlinear internal waves in the South China Sea, *Geophys. Res. Lett.*, **32**, L05615, doi:10.1029/2004GL022012, 2005.

Moore, S. E., and R.-C. Lien, Pilot whales follow internal solitary waves in the South China Sea, *Mar. Mamm. Sci.*, **23**, 1, 193–196, 2007.

Zhao, Z., V. Klemas, Q. Zheng, and X.-H. Yang, Remote sensing evidence for baroclinic tide origin of internal solitary waves in the northeastern South China Sea, *Geophys. Res. Lett.*, **31**, L06302, doi:10.1029/2003GL019077, 2004.

PUBLICATIONS

Lien, R.-C., T. Y. Tang, M. H. Chang, E. A. D'Asaro, Energy of nonlinear internal waves in the South China Sea, *Geophys. Res. Lett.*, **32**, L05615, doi:10.1029/2004GL022012, 2005.

Chang, M.-H., R.-C. Lien, T. Y. Tang, E. A. D'Asaro, and Y. J. Yang, Energy flux of nonlinear internal waves in northern South China Sea, *Geophys. Res. Lett.*, **33**, L03607, doi:10.1029/2005GL025196, 2006.

Moore, S. E., and R.-C. Lien, Pilot whales follow internal solitary waves in the South China Sea, *Mar. Mamm. Sci.*, **23**, 1, 193–196, 2007.

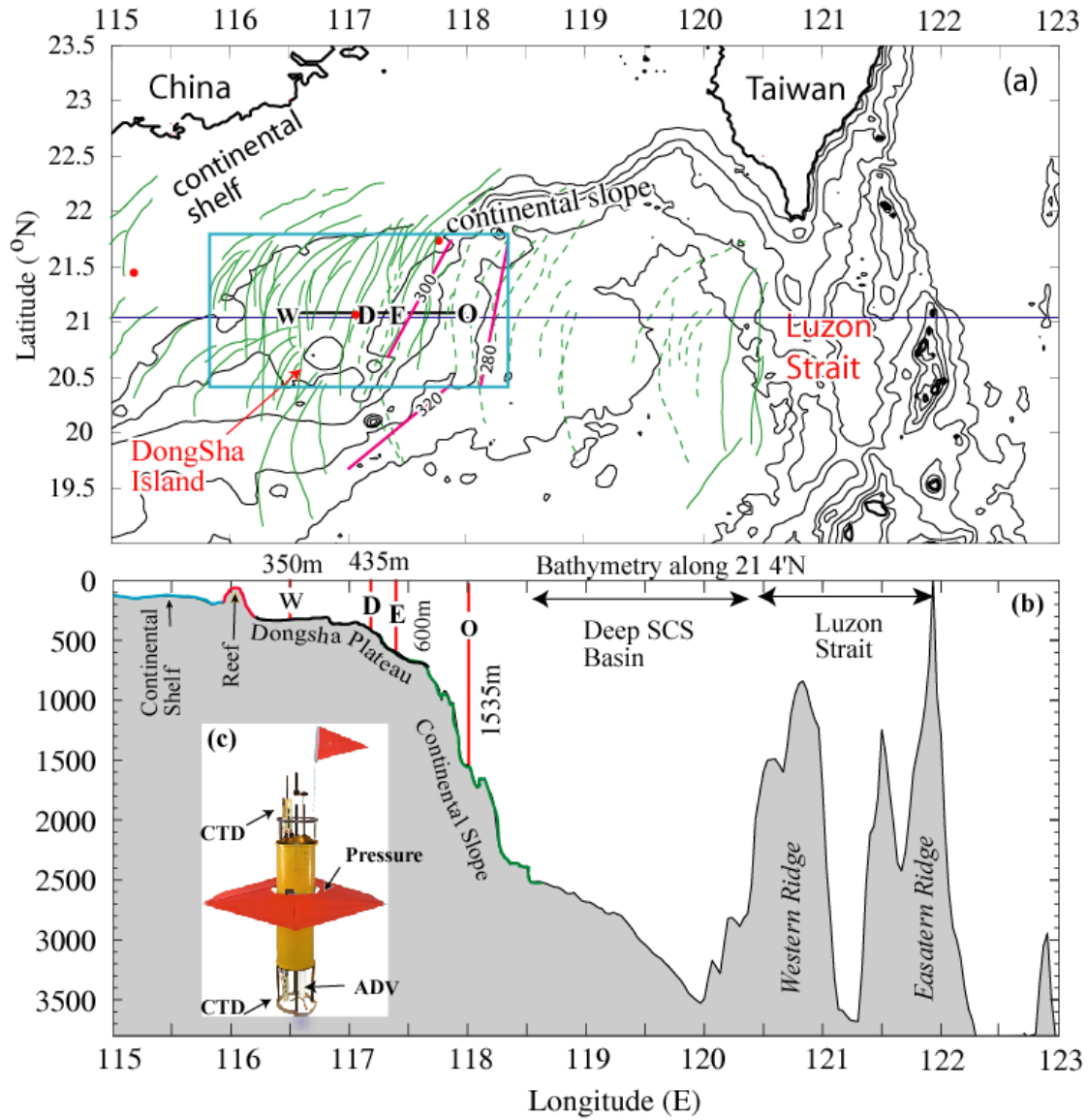


Figure 1: (a) Map of the northern South China Sea, and (b) bottom bathymetry along 21°4'N. In panel (a), green curves represent surface signatures of NLIWs identified in satellite images (Zhao et al., 2004), dashed for single-depression waves and solid for multiple wave packets. The blue box delineates the area where multiple wave packets are mostly found. Four primary stations in our April-May 2005 cruise are labeled as O, E, D and W. Shipboard and float measurements are taken along O-E-D-W. Three magenta curves illustrate isobaric orientations on the continental slope. In panel (b), two submarine ridges in the Luzon Strait are labeled. They are responsible for generating strong internal tides. Depths at four primary stations are also labeled. The inset (c) shows the Lagrangian float and sensors equipped on the float.

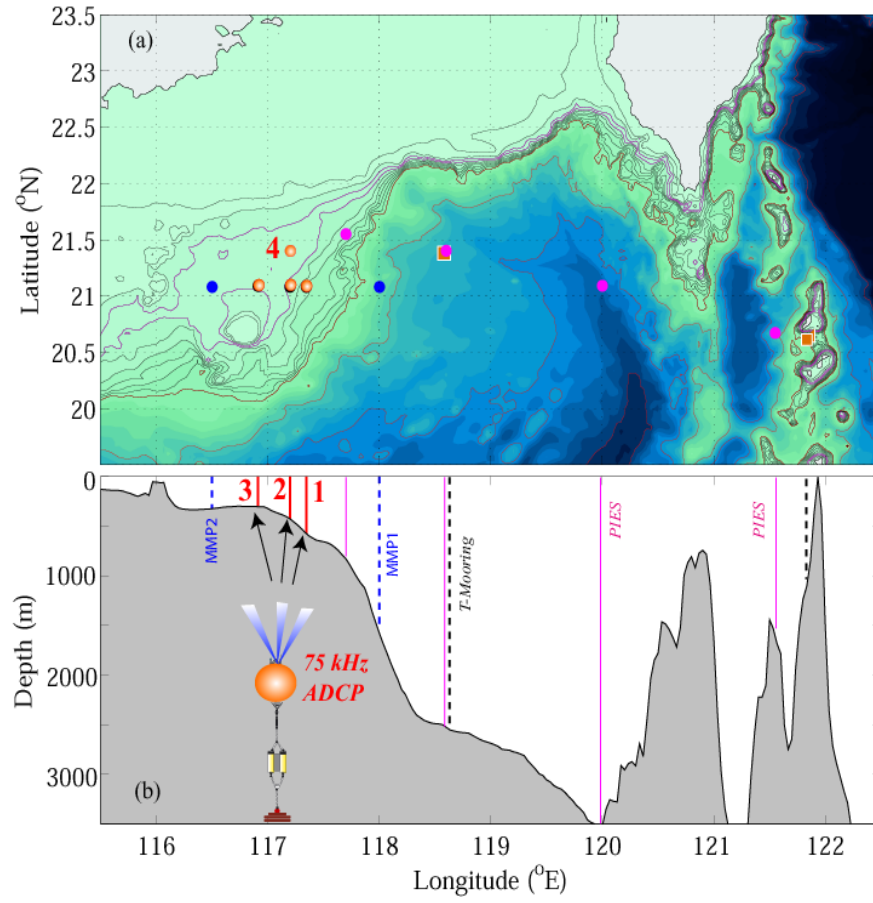


Figure 2: (a) Map of South China Sea and (b) bathymetry along 21°N. Four yellow bullets and red lines mark the locations of moored ADCPs. The configuration of the bottom mounted 75-kHz ADCP is shown in (b). Two blue dots (vertical blue dashed lines), magenta dots (vertical magenta lines), and brown squares (vertical black dashed lines) mark the positions of McLane moored profilers (Alford), PIES (Farmer), and temperature moorings (Tang and Ramp), respectively. Labels 1-4 represent the moored ADCPs.

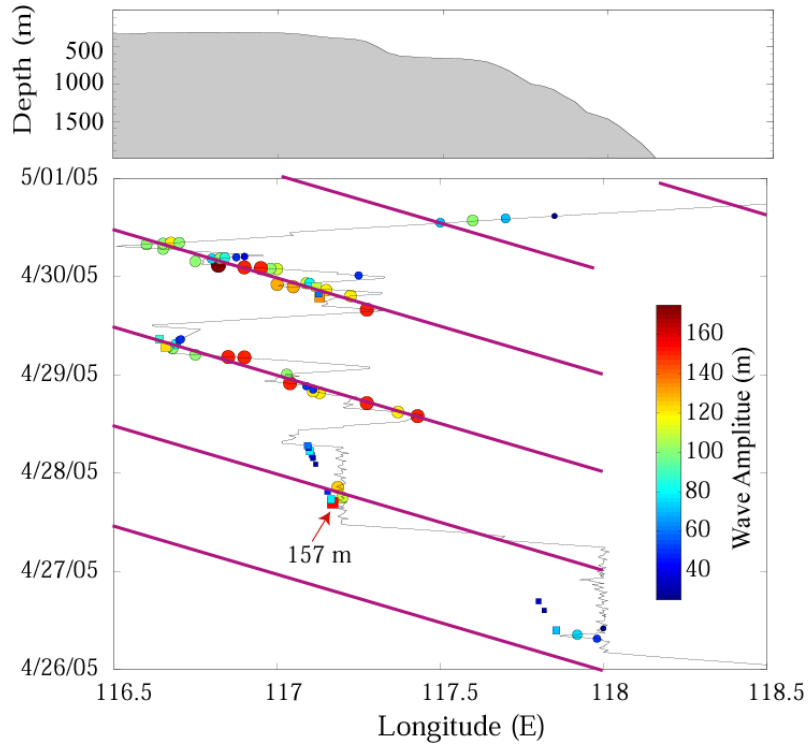


Figure 3: Compilation of amplitudes of NLIWs recorded on EK500 (color dots), shipboard ADCP, and Lagrangian float (filled squares). Colors and sizes of symbols represent amplitudes of NLIWs. The black curve indicates the ship track. Magenta curves illustrate the constant phase lines of observed NLIWs. NLIWs arrive diurnally.

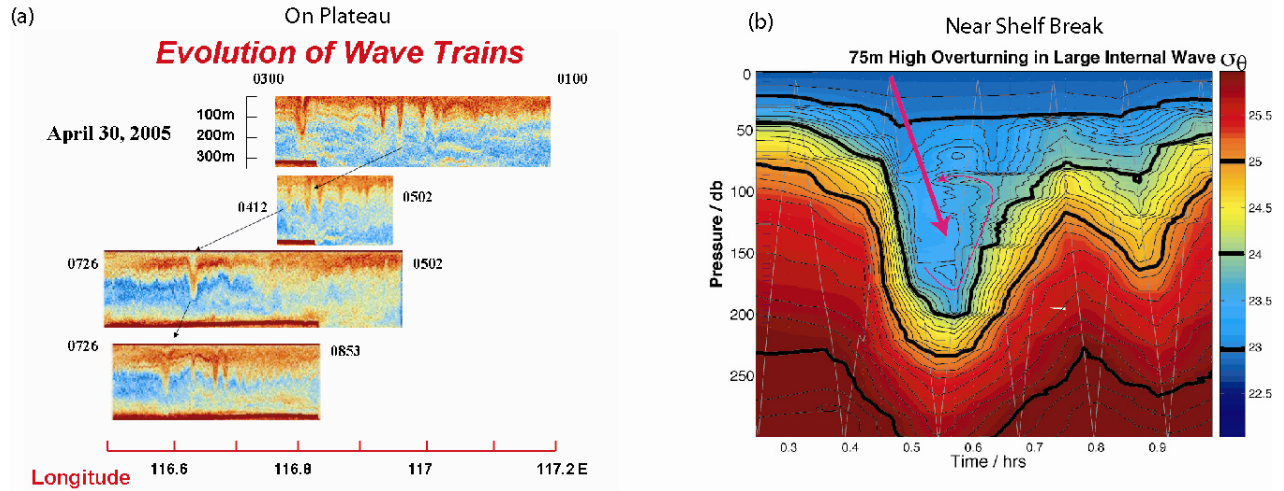


Figure 4: (a) Four sequential sections of EK500 images taken north of Dongsha Island illustrating the evolution of trains of NLIWs, and (b) contour of isopycnal surfaces taken with yoyo CTD profiles illustrating the large-amplitude NLIWs and large turbulence overturning. These observations were taken onboard of Taiwanese R/V Ocean Research 3 in April of 2005. In panel (a) the beginning time and the end time of each section of EK500 image are labeled.

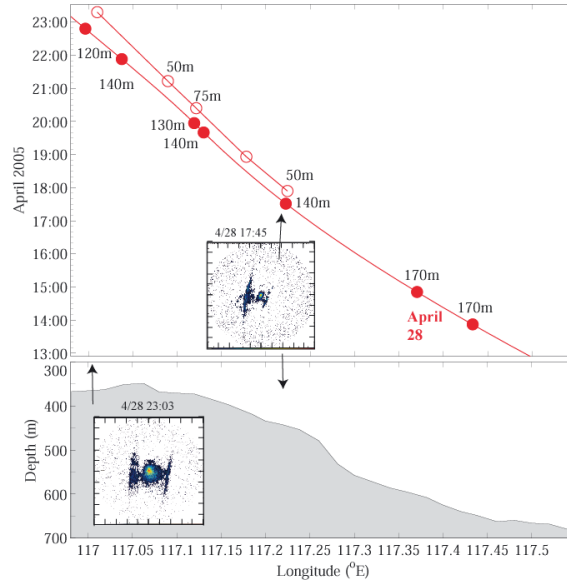


Figure 5: Tracking the evolution of a large-amplitude NLIW. Solid red dots in the upper panel show the locations and time of the tracked NLIW and open circles show the trailing wave, presumably developed from the tracked NLIW. Vertical amplitudes of NLIWs are labeled. Two insets are shipboard marine radar images showing the presence of two waves.

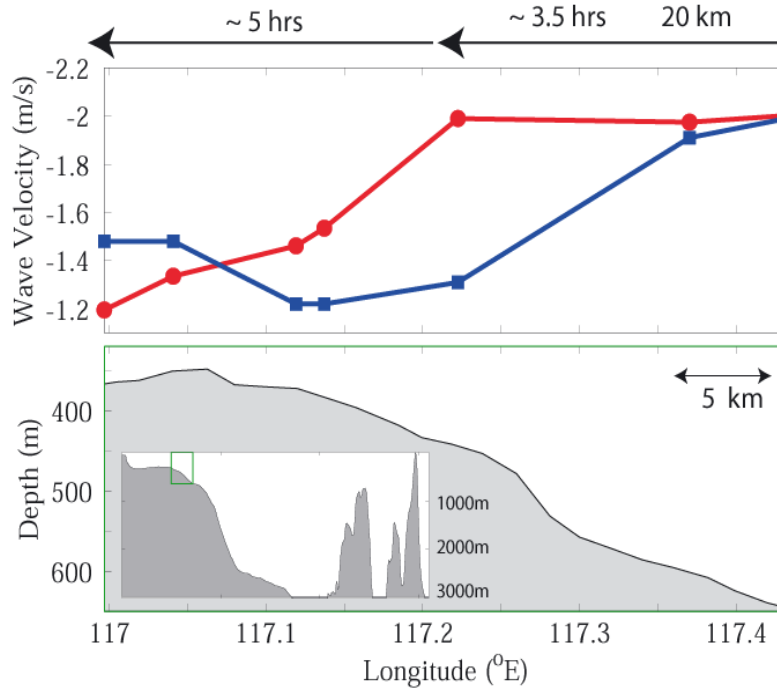


Figure 6: Evolution of maximum current velocity and propagation speed of NLIW on a shoaling topography. The upper panel shows the maximum westward zonal velocity (red curve) and the propagating speed of NLIW (blue curve). The bottom panel shows the bathymetry. The inset is the bathymetry along 21°N in the SCS. A spatial scale of 5 km is shown.

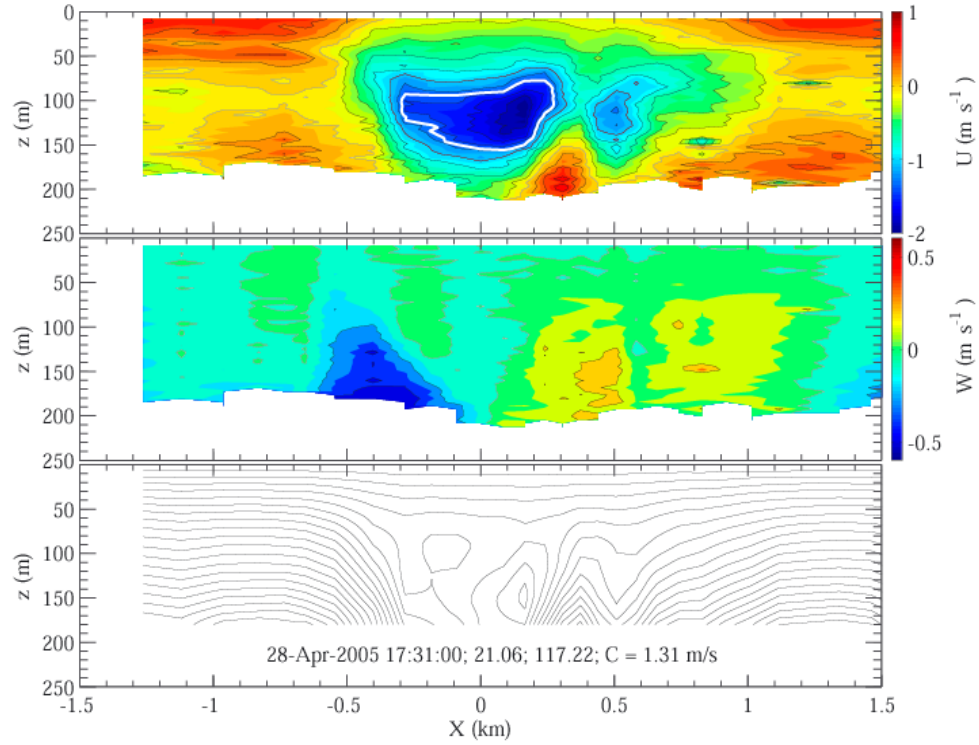


Figure 7: Illustration of the recirculation trapped core. The top two panels are contour plots of zonal velocity and vertical velocity observed by the shipboard ADCP during the 3rd encounter of the NLIW (Figs. 5 and 6). The bottom panel shows the streamline computed from the zonal velocity. The white curve in the top panel marks the contour line of the westward propagation speed of the NLIW.

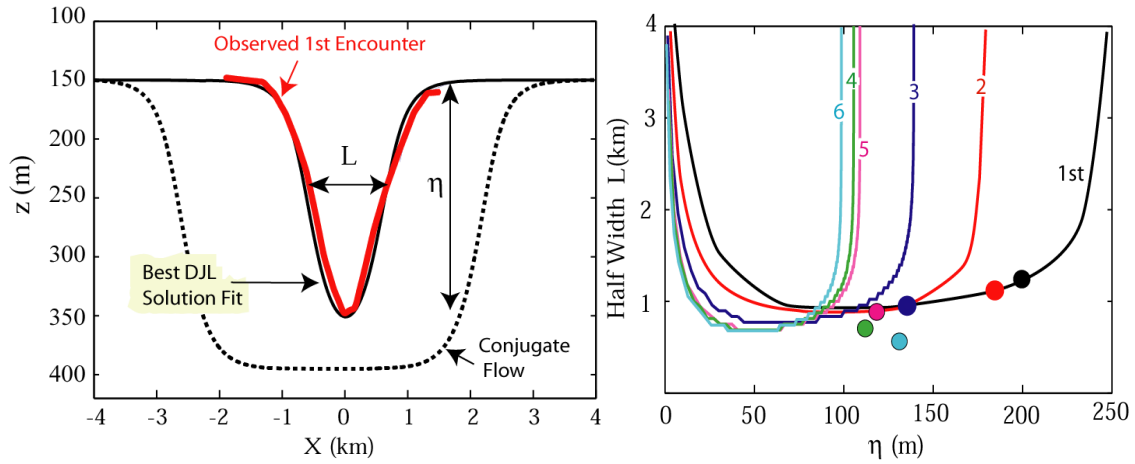


Figure 8: Observations compared with solutions of the model of Dubreil-Jacotin-Long (DJL). The left panel shows the good agreement between observations (red curve) at the first encounter with DJL (solid black curve). The right panel shows the comparison for observations in the first 6 encounters. Solid dots are observations and curves are DJL solutions. To agree with DJL solutions, solid dots should fall on curves of the same color.

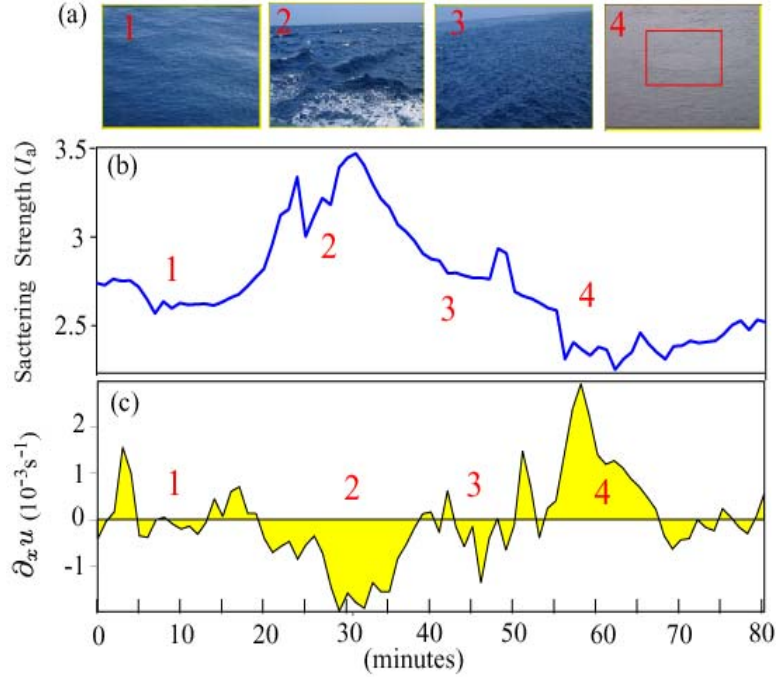


Figure 9: (a) Pictures taken during the experiment illustrating the surface signatures accompanying the passing NLIW, (b) the marine radar scattering strength, and (c) the horizontal velocity convergence computed from shipboard ADCP measurements in 8.4–12.4-m depth.

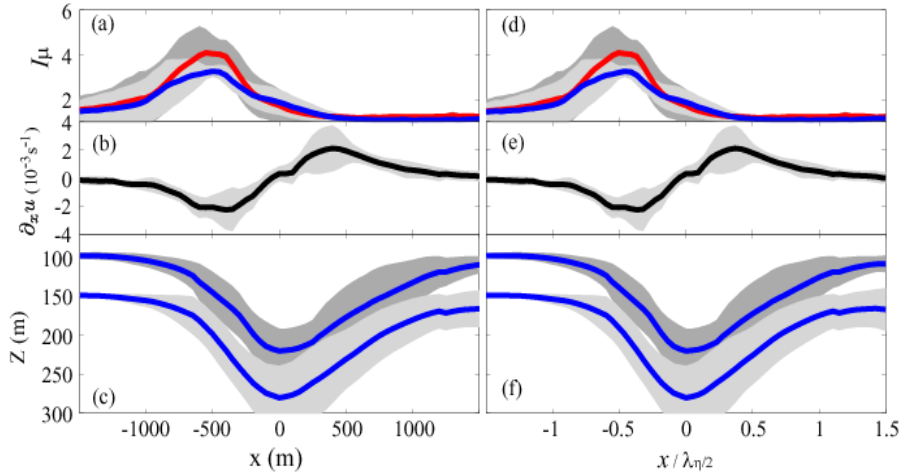


Figure 10: (a) Surface scattering strength, (b) horizontal velocity convergence, and (c) vertical displacement of NLIWs at initial depths of 100 and 150 m averaged over seven NLIW events. (d)–(f) are the same as (a)–(c) except that the x-axis is scaled by the wave width of the half maximum amplitude. Red curves in (a) and (d) represent the surface scattering strength observed from ahead of the propagating NLIWs and the blue curves represent the surface scattering strength observed from behind the propagating NLIWs. The light gray and heavy gray shadings represent 95% confidence intervals.

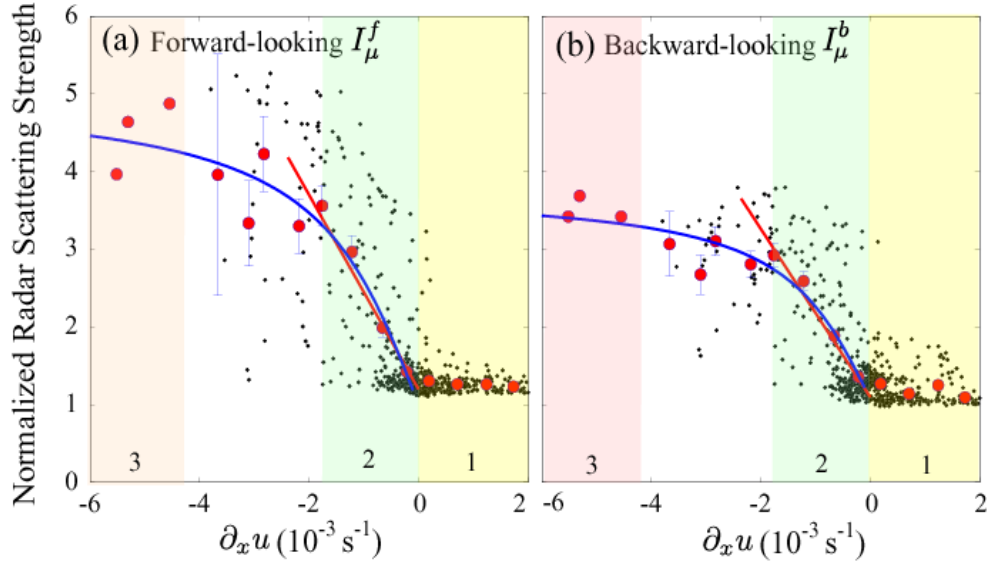


Figure 11: Scatter plots and model fits between the surface scattering strength and the horizontal velocity convergence computed from shipboard ADCP measurements. (a) Observations (gray dots) taken ahead of the propagating NLIWs. Red dots represent averages over constant intervals of horizontal convergence 0.005 s^{-1} . Vertical lines represent 95% confidence intervals. Red and blue curves represent the linear and arctangent fits to observations (red dots). Three regions labeled as 1, 2, and 3 represent the divergence zone, the weak convergence zone, and the strong convergence zone, respectively. Symbols, curves, and labels in the panel (b) are the same as in (a), but observations were taken from behind the propagating NLIWs.

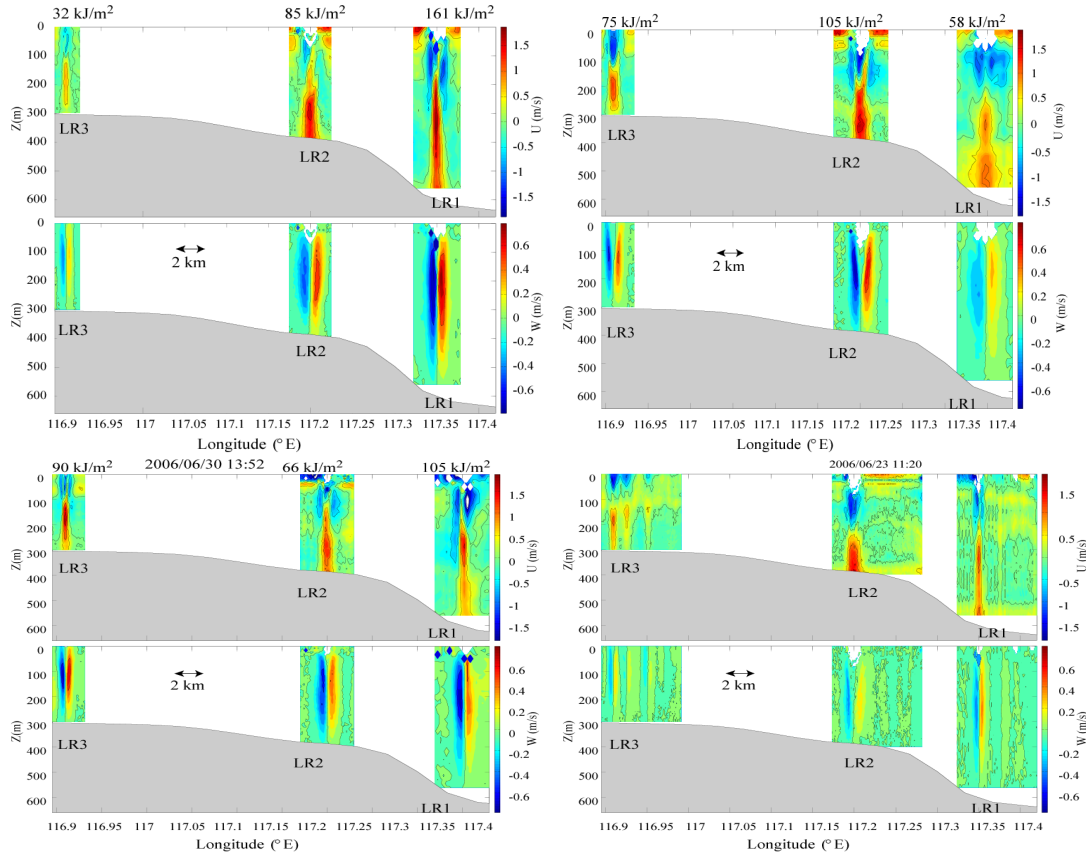


Figure 12: Illustration of the energy evolution of shoaling NLIWs revealed by long-term moored ADCP measurements. NLIWs propagate prevailingly westward. The depth integrated total kinetic energy is shown at the top of panels.

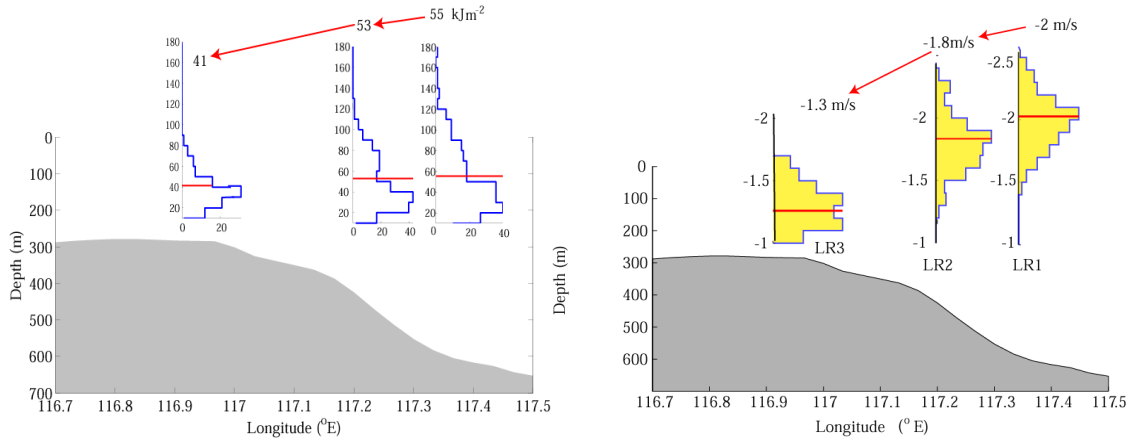


Figure 13: Histograms of the total kinetic energy (left panel) and propagation speed of NLIWs estimated at three mooring sites. Red horizontal lines represent the mean values.

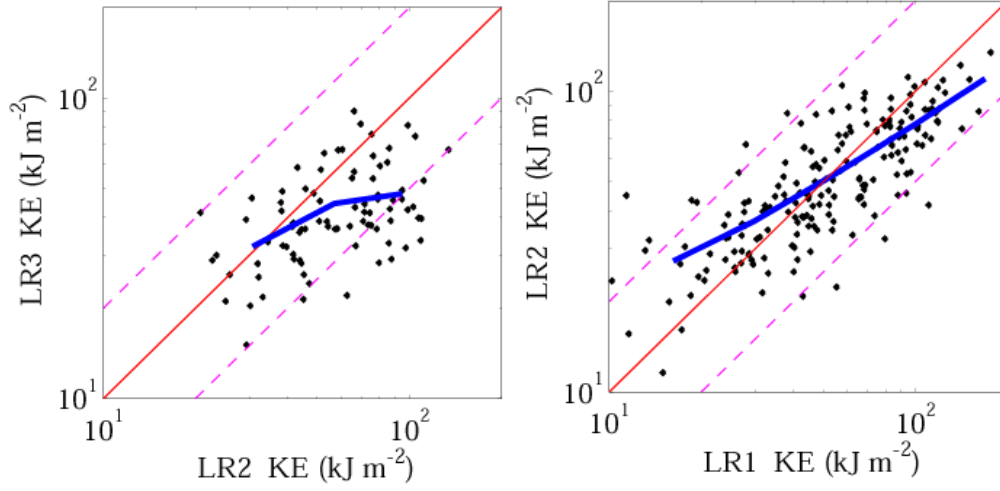


Figure 14: Comparison of the total kinetic energy of NLIWs at three mooring sites. The total kinetic energy is computed after removing the background current (dots). Blue curves are bin averages of black dots.

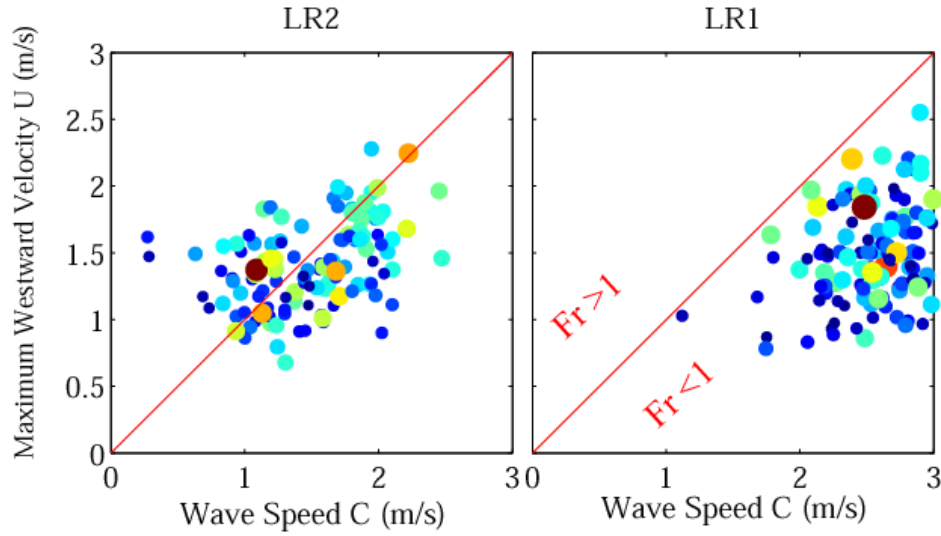


Figure 15: Comparison between the maximum westward velocity and the propagation speed of NLIWs. Colors and sizes of dots represent magnitudes of the total kinetic energy of NLIWs. Red lines indicate $Fr = U_{max}/C = 1$. Trapped cores exist in the region of $Fr > 1$.

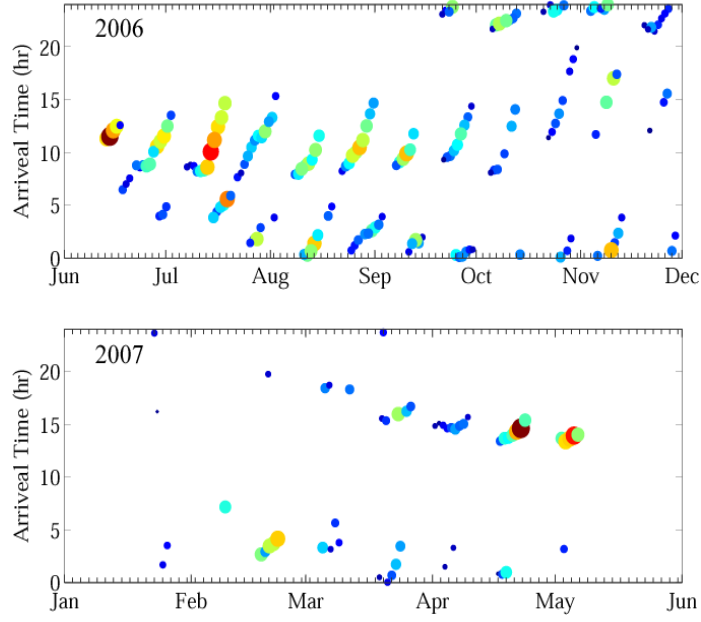


Figure 16: Arrival time of NLIWs in local hour of day at LR1. Sizes and colors of symbols represent the total kinetic energy of NLIWs.

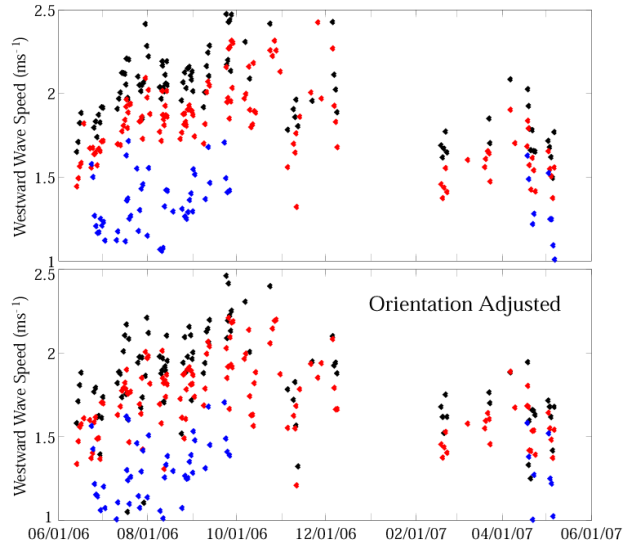


Figure 17: The westward propagation speed of NLIWs (upper panel) and the propagation speed in the wave direction (lower panel). Black, red, and blue dots are estimates of propagation speeds at LR1, LR2, and LR3, respectively.

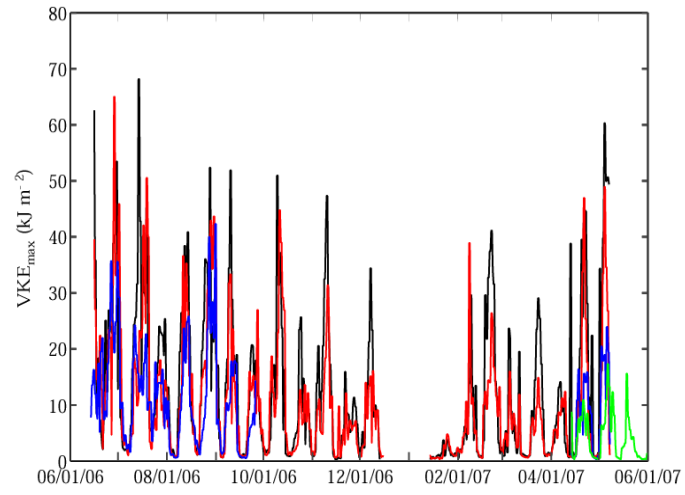


Figure 18: Depth integrated vertical kinetic energy of NLIW. Black, red, and blue curves represent estimates at LR1, LR2, and LR3, respectively.